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Predicting hydrogen and oxygen indices (HI, OI) from conventional well logs using a Random Forest machine learning algorithm

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ABSTRACT

Hydrogen Index (HI) and Oxygen Index (OI) are two critical parameters for assessing the hydrocarbon potential and depositional environment of any source rocks. The most common method to measure these values is to use programmed pyrolysis on drill samples. However, this method can be very time consuming, expensive, and in many cases much of the well bore may be overlooked due to biased sampling. Geochemical parameter predictions from wireline logs (i.e., Passey) have been used in the past to varying success. This is largely because petrophysical predictions often attempt to solve for linear regression solutions where this may not be the case. Here we evaluate the use of a Random Forest (RF) machine learning (ML) model to predict HI and OI from four wells from the offshore east coast of Newfoundland, Canada. The model was trained and tested using programmed pyrolysis data, organic petrology techniques, and wireline logs for prediction. The model was evaluated using mean absolute error (MAE), root mean square error (RMSE), correlation of determination (R^2), and Spearman's rank correlation (R2). Excellent correlation coefficients were observed for RF model predictions for HI and OI that range 0.90 to 0.98 and 0.90 to 0.95 R^2 respectively. The MAE for HI and OI values range 17.30 to 52.48 and 2.82 to 12.79 respectively. The RMSE for HI and OI range 21.43 to 71.51 and 3.85 to 16.82 respectively. The Spearman's rank correlation for HI and OI range 0.87 to 0.97 and 0.90 to 0.96 respectively. This study confirms that the use of ML models can be extremely useful to predict geochemical parameter from wireline logs.

1. Introduction

Source rock evaluation in conventional hydrocarbon systems begins by assessing the quality, quantity, and thermal maturity of organic matter [\(Carvajal-Ortiz and Gentzis, 2015](#page-24-0)). A typical first pass approach to assess this is by using programmed pyrolysis in an attempt to understand hydrocarbon generation and expulsion and eventual accumulation and production. A good source rock would be one that is considered to have a high total organic carbon (TOC) content, however not all organic carbon has the potential to generate hydrocarbon [\(Tissot](#page-24-0) [et al., 1974](#page-24-0)). Organic matter must be associated with hydrogen to be able to generate significant hydrocarbon. Therefore, a source rock with high hydrogen content is desirable. Hydrogen content is estimated by programmed pyrolysis by measuring the amount of hydrocarbons formed during the thermal decomposition of organic matter in the sample. This is measured in milligrams of hydrocarbons per gram of rock and is noted in programmed pyrolysis as S2 [\(Espitalie et al., 1977](#page-24-0)).

Hydrogen index (HI) is equal to (S2/TOC) x 100. Oxygen index (OI) which is a measure of S3 from programmed pyrolysis in milligrams of CO2 per gram of rock is equal to (S3/TOC)x100. High OI in a source rock is indicative of gas-prone, terrigenous-sourced, kerogen type III and IV source rock and hence is considered undesirable quality for a good oilprone source rock. The programmed pyrolysis method (e.g., Rock-Eval, HAWK, etc.) is a good first approach to evaluate hydrocarbon potential based on the aforementioned criterial, but one must use extreme caution in using programmed pyrolysis as a stand-alone dataset (e.g., [Dembicki, 2009](#page-24-0)). Other datasets must be integrated in order to understand the sedimentary systems that deposited the source rock and understand the type, distribution, and preservation of the organic matter.

For many years researchers have attempted with variable success to relate geochemistry data to wireline log data in an attempt to relate these geochemical parameters to well log information to predict good source rock intervals (e.g., [Passey et al., 1990, 2012](#page-24-0); [Creaney and](#page-24-0)

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Oil field Bathymetry Gas field

Fig. 1. . Area map. a.) Map of Canada showing the location of Newfoundland and the study area. b.) Map showing the location of the Jeanne d'Arc, Flemish Pass Basins and the Central Ridge. c.) Map showing the location of the three wells in this study from the Central Ridge (yellow dots). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Fig. 2. Jeanne d'Arc Basin lithostratigraphic chart showing the Tithonian and Kimmeridgian intervals of interest. The Rankin Formation and the Egret Member are highlighted (Modified from [Enachescu, 2005](#page-24-0)).

[Passey, 1993\)](#page-24-0). However, due to the complexity of comparing geochemistry data and well log data, many of the relationships are nonlinear and fail to give accurate results when using only linear regression models [\(Wang et al., 2019a, 2019b](#page-24-0)). A relatively new approach to this problem is using machine learning algorithms than can produce more accurate results than traditional statistical analysis and can handle large datasets producing fast and cost-effective results [\(Rodriguez-Galiano](#page-24-0) [et al., 2015\)](#page-24-0).

The aim of this study is to test the predictive accuracy of Random Forest (RF) analysis, a machine learning algorithm based on an ensemble of decision trees ([Breiman, 2001](#page-24-0)) to predict HI and OI using only wireline logs as input data on four east coast wells from offshore Newfoundland. Random Forest analysis has a proven track record in many scientific and engineering fields and has proven predictive accuracy in classification and regression analysis [\(Handhal et al., 2020](#page-24-0)). First, programmed pyrolysis, and organic petrology techniques will be used to gain an understanding of source rock types, organic matter types, and the comparison to HI and OI parameters. Second, run a machine learning Random Forest analysis using a normalized triple combo suite of well logs from the four east coast wells to predict HI and OI, Third, compare the results to the known source rock and organic matter types to test the Random Forest prediction.

2. Geological background

This study is based on three wells drilled in the Central Ridge area located off the east coast of Newfoundland approximately 400 km east of St. John's [\(Fig. 1](#page-1-0)). The fourth well's name and location cannot be disclosed due to the proprietary nature of the data and will be referred to simply as Well A in this study. The Central Ridge is located in the Flemish Pass sub-basin. The Flemish Pass sub-basin formed in response to Late Triassic to Paleocene rifting events that formed deep sedimentary basins bounded by faults ([Creaney and Allison, 1987; Enachescu, 2005](#page-24-0)). The three Central Ridge wells were drilled conventionally between 1980 and 1988 in search of hydrocarbon bearing Jurassic and/or Cretaceous

sandstone intervals with little success ([Cotterill, 1987](#page-24-0), unpublished). In the Central Ridge area, the depositional environment has been interpreted as deltaic to prodelta deposits [\(BeicipFranlab, 2015](#page-23-0)). The reservoir sandstone intervals are sandwiched in between thick intervals of organic-rich and/or organic-lean mudrocks [\(BeicipFranlab, 2015\)](#page-23-0). The mudrock intervals are the focus of this study.

The Kimmeridgian-aged Egret Member of the Rankin Formation (Fig. 2) in offshore Newfoundland has been considered the primary source rock in the Jeanne d'Arc Basin and has been extensively studied in the past [\(Swift and Williams, 1980](#page-24-0); [Creaney and Allison, 1987; Fowler](#page-24-0) [et al., 1990, 1991;](#page-24-0) [Huang, 1994](#page-24-0); [Fowler and McAlpine, 1995](#page-24-0); [DeSilva,](#page-24-0) [1999; Enachescu et al., 2010](#page-24-0), [Enachescu, 2012\)](#page-24-0). It has been described as an excellent hydrocarbon potential marine source rock ([Swift and Wil](#page-24-0)[liams, 1980](#page-24-0)) and is considered to be the equivalent source rock found not only in Kimmeridgian sediments, but also preserved in Tithonianaged sediments throughout other subbasins in offshore Newfoundland including the Central Ridge with high hydrogen index, high TOC content, and low oxygen index [\(Enachescu, 2005](#page-24-0); [Fowler et al., 2007](#page-24-0)). The depositional environment of this source rock has been interpreted as pelagic due to the observed lack of silt and sand grains and abundance of well preserved marine algal organic matter [\(Raine, 2006,](#page-24-0) unpublished). Other fine grained sediments with high organic content are also present in the thick mudrock intervals but have very little hydrocarbon potential due to the abundance of allochthonous continental derived degraded and reworked organic matter suggestive of a more terrigenous derived sediment likely from a deltaic source [\(Raine, 2006,](#page-24-0) unpublished). These mudrocks have low HI and high OI values with variable TOC content.

3. Sampling, data used and methodology

3.1. Programmed pyrolysis

A total of twenty-nine side wall core (SWC) samples and ninety drill cuttings samples were analyzed for hydrocarbon potential using programmed pyrolysis. The samples were sent to the Lithospheric Organic Carbon (LOC) laboratory, Department of Geoscience, Aarhus University in Denmark, for HAWK pyrolysis analysis for the standard cycle of Rock Eval 6 analysis. The samples were finely ground, and the method was carried out as described by [Lafargue et al. \(1998\)](#page-24-0). In the analyzer, the samples are heated to an iso-temperature of 300 ◦C and held for 3 min followed by a ramping of the temperature by 25 $°C/min$ to a temperature of 650 ◦C. S1 (free hydrocarbon) and S2 (thermally cracked kerogen) are the output parameter peaks reported in mg HC/g. The oxygen containing carbon in the kerogen released during the heating process produces the output parameter S3 and is reported as mg $CO₂/g$. The analyzer is then heated to a final temperature of 850 ◦C to oxidize the residual organic carbon and is reported as the total organic carbon (TOC: wt%) of the sample [\(Lafargue et al., 1998](#page-24-0)).

3.2. Organic petrology

Four SWC and nineteen drill cuttings samples were selected for organic petrology examination based on HI and OI variations from the pyrolysis data. The samples were prepared into finely polished epoxyresin sample pellets. A Zeiss Axioimager II microscope equipped with the Diskus-Fossil system was used to carry out all of the organic petrology analysis at the GSC in Calgary, Canada. No less that 50 vitrinite reflectance measurements were taken on each sample at 50×

Fig. 3. Random Forest flow chart and software programming used in this study (Modified from [Sun et al., 2019\)](#page-24-0).

Table 1

List of training data for Random Forest model.

Table 1 (*continued*)

Table 1 (*continued*)

$_{\rm GR}$	RESD	DENS	NPHI	DTC
API	ohm/m	Kg/m^3	v/v	μ s/m
72.52	22.42	2557.71	0.33	264.54
65.75	14.89	2508.05	0.30	281.50
77.79	6.70	2698.35	0.21	233.31
67.03	28.18	2549.06	0.31	286.40
46.61	53.97	2647.19	0.19	230.58
74.81	72.87	2424.40	0.32	350.96
81.14	7.48	2588.63	0.21	279.92
59.22	6.97	2702.25	0.13	202.35
98.48	4.90	2678.53	0.23	226.79
56.22	11.60	2719.31	0.08	191.44
33.74	1.50	2091.60	0.28	304.76
50.82	1.79	2112.40	0.42	262.96
45.90	2.60	2072.80	0.48	262.24
52.22	1.53	2958.70	0.28	297.84
47.82	2.51	2928.30	0.32	270.56
26.58	2.74	2728.20	0.51	369.24
53.87	3.12	2597.10	0.23	240.20
55.76	9.89	2622.20	0.24	243.96
65.55	3.46	2353.70	0.37	291.72
76.25	3.40	2409.80	0.35	295.00
74.96	3.68	2520.00	0.30	319.12
73.46	4.42	2557.10	0.35	289.20

Table 2

List of RF model parameters used.

$bootstrap = True$, ccp alpha = 0.0, $criterion = 'mse'.$ max depth $=$ None, max features $=$ 'auto', max leaf nodes $=$ None,	min samples leaf $= 1$, min samples split $= 2$, min weight fraction leaf $= 0.0$, n estimators = 25, n jobs = None, $oob score = False$,	
max samples $=$ None,	random state = 42 , verbose = 0,	
min impurity decrease $= 0.0$,	warm start $=$ False	
min impurity split $=$ None,		

magnification. An ultrafine measurement probe $(0.3 \mu m^2$ spot size) was used under oil immersion (refractive index, $n = 1.518$ at 23 °C). An yttrium-aluminum-garnet reference standard was used with a reflectance of 0.906% under oil immersion. Maceral point counts were carried out using a twenty-one cross-hair grid (e.g., [Gordon et al., 2021](#page-24-0)). No less than two-hundred maceral counts per sample were counted to produce an organic maceral distribution normalized to 100% of the measured TOC. No macerals were counted that appeared to be isolated in the sample binder. The maceral categories (vitrinite, inertinite, liptinite, and solid bitumen) were determined based on maceral attributes described in [ICCP \(1998\),](#page-24-0) [ICCP \(2001\),](#page-24-0) [Pickel et al., 2017,](#page-24-0) and [Sanei,](#page-24-0) [2020.](#page-24-0)

3.3. Well log data

Many studies have shown that wireline logs can be sensitive to the presence of organic matter in rocks [\(Passey et al., 1990;](#page-24-0) [Creaney and](#page-24-0) [Passey, 1993](#page-24-0); [Passey et al., 2012](#page-24-0); [Bolandi et al., 2017](#page-24-0); [Wang et al.,](#page-24-0) [2019a, 2019b](#page-24-0); [Handhal et al., 2020\)](#page-24-0). A single log parameter may be sensitive to certain downhole conditions and lithology, therefore multiple log parameters are required for RF prediction models ([Wang et al.,](#page-24-0)

[2019a, 2019b](#page-24-0)). The most common logs showing sensitivity to organic matter include the natural gamma, resistivity, transit interval time, and porosity (density and neutron). In general, the higher the organic content in the rock, the more obvious the response is to the wireline logs ([Wang et al., 2019a, 2019b\)](#page-24-0). The wireline log input data used for RF training in this study include gamma ray (GR), resistivity (RESD), density (DENS), neutron (NPHI), and sonic DTC). All wireline logs were mathematically normalized to each other using the histogram module in Interactive Petrophysics® software to improve data consistency and integrity in order to create a common basis for log comparison.

3.4. Random Forest and well log data

3.4.1. Random Forest modelling

Random Forest (RF) is a machine learning method that consists of an ensemble of randomized classification and regression trees (CART) that generates many decision trees to improve the performance of the prediction model [\(Breiman, 2001\)](#page-24-0). A decision tree represents a set of limits that are hierarchically organized and randomly applied from a root node as many times as the number of trees in the ensemble ([Keykhay-Hos](#page-24-0)[seinpoor et al., 2020\)](#page-24-0). The trees are generated from a subset of training samples through replacement (a bagging approach) and the same sample can be selected several times or may not be selected at all. Approximately two thirds of the samples (referred to as in-bag samples) are used to train the trees and the remaining one third (referred to as out-of-the bag samples) are used to confirm how well the RF model performs ([Breiman, 2001\)](#page-24-0). The final classification decision, or a committee vote, is taken by averaging (using the arithmetic mean) the class assignment probabilities calculated by all produced trees (Belgiu and Drăgut, 2016). The main advantage of using this approach is the RF classification algorithm can model non-linear relationships and the model consists of numerous random decisions trees. Each individual tree creates an uncorrelated forest of trees whose prediction by voting committee is more accurate [\(Grimm et al., 2008\)](#page-24-0). The RF workflow and software used in Programmed pyrolysis results.

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mgHC/gTOC $_{\text{mgCO}_2/\text{gTOC}}$

 172 33

Table 3 (*continued*)

Table 3 (*continued*)

this study is shown in [Fig. 3](#page-3-0). [Table 1](#page-4-0) shows all the training data from programmed pyrolysis and wireline log data used as input data.

3.4.2. RF feature selection and model evaluation

In this study a randomly selected twenty-five trees were used, and the input matrix (training data) are the well log curves used for each well (gamma ray, resistivity, density, neutron, and sonic). The default setting for the RF prediction model is shown in [Table 2.](#page-7-0) In previous studies the software parameters used can be set to adjustable user settings, however it is recommended to simply use software default settings as a first approach. Changing the model settings to higher randomness have indicated noise variables in final prediction data (e.g., [Svetnik et al.,](#page-24-0) [2003;](#page-24-0) [Grimm et al., 2008\)](#page-24-0). However, it should be noted that hyperparameter tuning can be applied in an attempt to improve the RF performance. [Probst et al., 2019](#page-24-0) have provided the tuneRanger R package that aids in tuning the RF model with model-based optimization (MBO). To evaluate the RF model, four statistical metrics were used: i) mean absolute error (MAE), ii) root mean squared error (RMSE), iii) the correlation of determination coefficient (R^2) , and iv) a Spearman's rank correlation (*Rs*) was used to determine the degree in which the data sets are correlated. A correction factor (c_f) was applied to Σd^2 when ranks were found to be tied. The correction factor was added to Σd^2 for each tied rank in the datasets. The MAE and RMSE are used to recognize the outliers in the dataset. R^2 is used to evaluate the accuracy of the model

where x_i and y_i are the measured and estimated values of HI and OI, \bar{x} and \bar{y} are their arithmetic mean, and n is the total number of measured HI and OI data points.

$$
MAE = \frac{\sum_{i=1}^{n} |y_i - x_i|}{n}
$$
 (1)

$$
RMSE = \sqrt{\sum_{i=1}^{n} \frac{(\hat{y}_i - y_i)^2}{n}}
$$
 (2)

 $\hat{\mathbf{a}}$

$$
R^{2} = \frac{\text{explained variation}}{\text{total variation}} = \left(\frac{\sum_{i=1}^{n} (x_{i} - \overline{x})(y_{i} - \overline{y})}{\sqrt{\sum_{i=1}^{n} (x_{i} - \overline{x})^{2} \sum_{i=1}^{n} (y_{i} - \overline{y})^{2}}}\right)^{2}
$$
(3)

$$
R_s = 1 - \left(\frac{6\Sigma d^2}{n^3 - n}\right) \tag{4}
$$

$$
c_f = \frac{m(m^2 - 1)}{12}
$$
 (5)

Fig. 4. Pseudo van Krevelen plot showing hydrogen index (HI) vs. oxygen index (OI) for all samples from the studied area.

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Table 4

ł,

Organic petrology data collected on selected samples.

Vitrinite reflectance

Table 4 (*continued*)

Fig. 5. Photomicrographs showing the different organic matter types. All of the photomicrographs were taken under incident-light with oil immersion and a 50 X objective. Photomicrographs a., c. e. and f. are taken under white light; photomicrographs b., and d. are taken under UV-fluorescence mode. The red scale bar represents 50 um in length. a.) Dark brown thick layers of lamalginite (Lam) from Well A. Large fragment of inertinite is also present (Int) b.) Same field of view as in a. but under fluorescence light showing bright green fluorescing algae. c.) Example from Panther P-52 showing only scattered algae fragments (Alg) with recycled vitrinite (RV) fragments and inertinite (Int) in a silty argillaceous matrix. d.) As in photo c. but under fluorescence light. e.) An example from South Tempest G-88 showing scattered recycled vitrinite (RV) and inertinite fragments (Int) is a silty argillaceous matrix. f.) Similar example from South Tempest G-88 showing recycled vitrinite (RV) and inertinite (Int) fragments. No marine algae are present in these samples. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Fig. 6. Plot showing the percentage volume distribution of liptinites versus vitrinite + inertinite, as measured by microscopic point counting. Increase in distribution of liptinites suggest change to the marine depositional environment.

3.4.3. Data loading

The rock data is compiled and hosted on a local computer drive in a Microsoft Excel file format and the well log data are stored as LAS files. The application software used for RF is Jupyter Notebook within Anaconda (Python version 3.8). All of the RF code are stored in Jupyter notebook as .jpynb format. The Python libraries used are shown in [Fig. 3.](#page-3-0) The Excel and LAS data is loaded with the panda read_excel function and the with the lasio read function into separate data frames. A class object is constructed for each well to hold information such as well name, LAS filename, and curves needed for training and stored in a dictionary. The model is evaluated in scikit_learn. All of the pyrolysis data below 400 ◦C Tmax were filtered out.

4. Results and discussion

4.1. Programed pyrolysis and organic petrology

All data collected from programmed pyrolysis is presented in [Table 3.](#page-8-0) Total organic carbon (TOC) content from all the samples collected $(n = 119)$ range from 0.28 to 14.96 (wt%). S2 vales range from 0.00 to 135.26 (mg HC/g). S1 values were not used in this study as various drilling fluids were utilized during drilling that can add contamination errors in measured S1 values. The HI and OI values range from 0 to 904 (mg HC/g TOC) and from 2 to 158 (mg $CO₂$ /g TOC), respectively. Tmax values range from 418 to 445 ◦C.

The pseudo van Krevelen plot [\(Fig. 4](#page-12-0)) for all the samples shows a wide range of OM types from Type I oil prone, Type II oil prone, Type II/ III mixed, and Type III gas prone. Pseudo van Krevelen diagrams typically show a trend for organic matter that shows a depletion of HI and OI along the defined kerogen type lines, irrespective of the organic matter type, due to thermal maturity of the organic matter ([Peters et al., 2015](#page-24-0)). The samples from the four wells in this study show a progressive depletion of HI combined with increase in OI that is in contrast to the typical trend. The OM preservation here is controlled by oscillation in sea level from a marine carbonate-rich influenced OM maceral composition to more siliciclastic-rich terrigenous littoral and/or deltaic

influenced OM composition. This OM composition oscillates from a liptinitic-rich and carbonate-rich aquatic depositional environment with abundant production and exceptional preservation of hydrogen-rich algal matter to a more oxygen-rich OM aquatic depositional environment typical of a terrigenous sediment input with recycled and reworked OM along with silt and sand-sized siliciclastic minerals ([Omura and Hoyanagi, 2004; Hackley et al., 2020; Gordon et al., 2021](#page-24-0)).

Organic petrology results show a wide variety of maceral types are present in all the wells and age intervals including vitrinite (Type III), high-reflecting reworked vitrinite (Type IV) exhibiting a brighter grey colour than vitrinite, liptinite (Type II), inertinite (Type IV) exhibiting bright grey color and high %VRo, and solid bitumen [\(Table 4](#page-13-0)). These data are consistent with the programmed pyrolysis results. Representative photomicrographs of the organic petrology are illustrated in [Fig. 5a](#page-14-0) to f.

Maceral point count data are shown in [Table 4](#page-13-0) and the data are normalized to measured TOC values. By plotting the sum of the relative volume of primary and reworked vitrinite plus inertinite $(V + I)$ versus the relative volume of liptinite macerals (L) two distinct end members are showing the terrigenous influenced depositional environment versus marine. These have a strong correlation coefficient ($R^2 = 0.95$) and the strong influence of depositional environment and the oscillations in relative sea level have on the composition, preservation, and distribution of OM (Fig. 6). These observations occur regardless of the depth or age of the sediment and appear in a cyclical pattern throughout the Tithonian and Kimmeridgian.

4.2. Random Forest analysis for source rock prediction

4.2.1. Data integration

To better aid the Random Forest algorithm to predict the best source rock intervals using the well logs based on measured HI and OI data, the samples described above can be divided into three simplified groups based on the pyrolysis data and organic petrology results. TOC, HI, and OI values define these groups as they either have i) excellent hydrocarbon potential, ii) transition, and iii) poor hydrocarbon potential ([Fig. 7\)](#page-16-0).

4.2.2. *Excellent hydrocarbon potential* $(n = 13)$

Pyrolysis data for this sample grouping shows HI ranging 622 to 904 (mg HC/g TOC), OI ranging 2.0 to 15.0 (mg $CO₂/g$ TOC), and TOC ranging 3.30 to 14.96 (wt%; [Fig. 7](#page-16-0)). These samples exhibit fine laminations consisting of layers of calcite and OM. Fine disseminated pyrite crystals are associated with OM. Silt and sand-sized siliciclastic grains are rare and minor amounts of calcispheres and forams are also present indicating fully marine depositional environment. Point count data $(n =$ 5) shows liptinite to be the main OM present (92.0 to 95.0%). Minor amounts of reworked vitrinite and inertinite (3.0 to 18.0%), and bituminite (0.0 to 2.0%) are also present. Fine layers of bright yellow-green algae composed of thin-walled colonial or unicellular algae occur as distinct laminae. The fluorescence color of the lamalginite in these samples is consistent with early oil window maturity. Representative photomicrographs of the point counted maceral types are illustrated in [Fig. 5a](#page-14-0) and b.

Fig. 7. Hydrogen index (HI) vs. oxygen index (OI) plot showing three groups defined by programmed pyrolysis: i) excellent hydrocarbon potential, ii) transition, and iii) poor hydrocarbon potential.

Well A Panther P-52 120 120 90 $\overline{9}$ $R^2 \in 0.00$ 100 100 80 80 ϵ $^{\prime}$ \sim α 70 70 80 $\overline{}$ 60 60 $\widetilde{\mathbb{G}}$ $R^2 = 0.04$ GR GR GR 60 $R^2 = 0.05$ $5[°]$ 60 50 $R^2 = 0.05$ 40 Λ ⁰ \blacksquare 30 30 \bullet 20 20 $\overline{20}$ $\overline{20}$ 10 10 \circ \circ $\,$ $\,$ 150 100 1000 100 Ol mgCO₂/g 200 400 800 5^o 200 400 600 800 \mathfrak{c} 50 150 200 600 HI mgHC/gTOC OI mgCO₂/g HI mgHC/gTOC 700 700 25 25 \odot \bullet $R^2 = 0.59$ 600 600 $\overline{20}$ 20 500 500 α **BSD** 15 ESD 400 읍 ⁴⁰⁰ 15 RESD $\ddot{}$ $R^2 = 0.17$ 300 300 10 10 200 \bullet 200 $R^2 = 0.01$ ó $R^2 = 0.05$ $10[°]$ 100 \sim $_{\odot}$ \overline{a} 400 600
HI mgHC/gTOC 800 1000 100 150 200 200 400 600 800 50 100 150 50 OI mgCO₂/g OI $mgCO₂/g$ HI mgHC/gTOC 2750 2750 3000 3000 2700 2700 **Bought** 2500 2500 2650 2650 2000 2000 2600 2600 $R^2 = 0.23$ **SEN** 2550 $R^2 = 0.09$ $\begin{array}{c}\n 2560 \\
2550\n \end{array}$ $\frac{2}{5}$ 1500 **SEN** 1500 2500 2500 1000 1000 $R^2 = 0.29$ 2450 2450 $R^2 = 0.09$ g fr 500 Sh. 500 2400 2400 2350 2350 200 400 200 \circ 600 800 50 150 ϵ 400 600 800 1000 50 100 150 200 \circ 100 HI mgHC/gTOC HI mgHC/gTOC OI mgCO₂/g $Cl mgCO₂/g$ 0.4 0.4 0.6 0.6 0.35 0.35 0.5 0.5 0.3 0.3 ó 0.4 0.4 \overline{a} 0.25 0.25 $\overline{0}$ $R^2 = 0.00$ $\frac{1}{2}$ 0.25
2 0.2 $\circ \cdot$ $\frac{1}{2}$ 0.25
 ≈ 0.2 $\overline{\Xi}$ _{0.3} $\begin{array}{l} \Xi \\ \Xi \\ 0.3 \end{array}$ $R^2 = 0.28$ ٩ 0.15 0.15 0.2 0.2 0.1 0.1 \bullet $\ddot{}$ $R^2 = 0.18$ 0.1 0.1 0.05 0.05 $R^2 = 0.09$ \circ ċ C 400
HI mgHC/gTOC 200 600 800 50 100 150 1000 400 600 800 100 150 200 200 ϵ 50 \circ Ol mgCO₂/g HI mgHC/gTOC OI mgCO₂/g 350 350 400 400 350 300 300 350 $R^2 = 0.00$ θ α 300 300 250 250 250 250 $\frac{10}{5}$ 200 ö $\frac{10}{5}$ 200 DTC DTC \bullet 200 200 $R^2 = 0.24$ 150 150 $R^2 = 0.21$ 150 150 $R^2 = 0.13$ 100 100 100 100 50 50 50 50 \circ $\overline{0}$ \circ Ω 50 $10[°]$ 150 800 400 600
HI mgHC/gTOC 100
Ol mgCO₂/g 200 400 200 800 1000 50 150 200 600 $O\vert$ mgCO₂/g HI mgHC/gTOC

Fig. 8. Matrix scatter plots of HI and OI with various wireline logs.

Fig. 8. (*continued*).

Panther P-52

South Merasheen K-55

Fig. 9. Scatter plots of each well showing excellent correlation of predicted Random Forest data and programmed pyrolysis data. Note that some data are not normally distributed. Spearman's rank correlation was applied, and all data was found to have very high correlations.

Table 5

Error matrices for both HI and OI.

4.2.3. Transition (n = *31)*

This sample grouping represents a transition from excellent hydrocarbon potential to poor ([Fig. 7](#page-16-0)). Pyrolysis data for show HI ranging 401 to 599 (mg HC/g TOC), OI ranging 7.0 to 24.0 (mg $CO₂/g$ TOC), and TOC ranging 1.38 to 8.62 (wt%). These samples exhibit laminations with abundant silt-sized siliciclastic gains with scattered OM in abundant calcareous and argillaceous matrix. Rare fossil fragments are also present. Point count data ($n = 5$) shows liptinite to still be the main OM present (68.0 to 97.0%). However, reworked vitrinite $+$ inertinite (3.0 to 27.0%), and bituminite (0.0 to 17.0%) are showing more abundance than group above [\(Fig. 5](#page-14-0)c and d).

4.2.4. Poor hydrocarbon potential $(n = 75)$

This sample grouping shows HI ranging 0 to 393 (mg HC/g TOC), OI ranging 11.0 to 158.0 (mg $CO₂/g$ TOC), and TOC ranging 0.28 to 5.08 (wt%; [Fig. 7](#page-16-0)). These samples exhibit massive to laminated calcareous to silty/sandy mudrocks to siltstones to OM-poor lime mudrocks with very low hydrocarbon potential. A wide range of preserved OM are present in these samples. Point count data $(n = 13)$ shows scattered liptinite (24.0) to 89.0%). An abundance of reworked vitrinite $+$ inertinite (11.0 to 66.0%), and bituminite (0.0 to 14.0%; [Fig. 5e](#page-14-0) and f).

4.2.5. Thermal maturity

The sample groupings above, being based solely on TOC, HI, and OI, ignores the thermal maturity of the OM. Measured VRo% on vitrinite macerals ranges 0.25 to 0.87%. Converting Tmax values to %VRo equivalent (Eq. 6; [Jarvie, 2012\)](#page-24-0) ranges 0.40 to 0.81% which is in good relation to the measured data [\(Table. 4](#page-13-0)). Caution should be used with this equation (Eq. 6) to convert Tmax to %VRo as there is no universal correlation for Tmax and %VRo and the equation likely differs from mudrock unit to mudrock unit globally [\(Yang and Horsfield, 2020\)](#page-24-0).

$$
\% VRo \; equivalent = (0.0180 \times \text{Tmax}) - 7.16 \tag{6}
$$

These data indicate immature to oil window thermal maturity. However, both of these thermal maturity methods ignore the contribution of algal OM as these maceral types cannot be measured using these techniques ([Thompson-Rizer and Woods, 1987](#page-24-0)). [Gordon et al., 2021](#page-24-0) showed by integrating Fluorescence Red/Green Quotient (R/G Q) spectral data collected on a subsample set used in this study, measured only on preserved fluorescing algae, the %VRo equivalent ranges 0.48 to 0.61% indication that the thermal maturity has not yet reached the primary oil generation window. The variation of these measured values is due to the mixing and dilution of the higher reflecting and reworked vitrinite macerals, sourced from different depositional environments caused by variation in sea level ([Gordon et al., 2021](#page-24-0)). Therefore, the R/G Q %VRo equivalent is likely a more accurate estimation of the thermal maturity in these samples.

4.3. RF model performance and prediction

Prior to the RF prediction model, the relationships between HI, OI, and the well log parameters were investigated using a series of scatter plots [\(Fig. 8\)](#page-17-0). The results show no correlation between the HI, OI and the well log parameters. HI and OI were the only predicted attributes used in the RF model.

The RF modelling is predicting HI and OI with significant correlation

coefficients that range from 0.90 to 0.98 and 0.90 to 0.95 R^2 respectively ([Fig. 9\)](#page-19-0). The MAE for HI and OI values range from 17.30 to 52.48 and 2.82 to 12.79, respectively. The RMSE for HI and OI range from 21.43 to 71.51 and 3.85 to 16.82, respectively. These values are relatively small given the overall large variation in HI and OI data. The Spearman's rank correlation for HI and OI range from 0.87 to 0.97 and 0.90 to 0.96, respectively (Table 5). The South Tempest G-88 has the lowest MAE and RMSE as the HI and OI variation is relatively small with HI ranging from 48 to 369 mg-HC/g TOC and OI ranging from 13 to 91 mg-CO₂/g TOC. Well A has the highest variation with HI ranging from 0 to 907 mg-HC/g TOC and OI ranging from 2 to 158 mg- $CO₂/g$ TOC. This indicates the error values is likely only significant as the HI and/or OI values approach the applied cut-offs mentioned above for each of the three simplified groupings. Therefore, one must use caution as HI and OI values approach preconceived cut-off values. It should be noted that data overfitting can be an issue when using complex machine learning algorithms. However, since each tree is trained on a unique subsample of the data using bootstrapping techniques, RF models are robust and resistant to overfitting [\(Grimm et al., 2008](#page-24-0); [Handhal et al., 2020\)](#page-24-0). A visual comparison of the final output of predicted HI and OI compared to the measured results can be seen in [Fig. 10](#page-21-0) for each well. A pseudo van Krevelen comparison of the measured vs. the predicted HI and OI for each well can be seen in [Fig. 11.](#page-23-0) The HI prediction curves can predict the organic richness of each well where there were no samples taken. Similarly, the predicted OI graphs are emphasizing the depth where it is organic lean. These predictions further suggest organic richness is directly related to water depth and changes in depositional environment that affect the factors controlling the accumulation and preservation of organic matter.

5. Conclusions

This study tested the validity of using a Random Forest (RF) machine learning algorithm to predict hydrogen index (HI) and oxygen index (OI) using wireline logs. The RF model was trained using programmed pyrolysis data (Rock-Eval 6 method) and detailed organic petrology. The predictions of HI and OI from the RF model was evaluated using four statistical error matrices. The results showed the RF prediction performed very well with acceptable mean absolute error (MAE), root mean square error (RMSE), high correlation of determination (R^2) and high Spearman's rank correlations (R_2) . This study showed that by using an integrated approach and using machine learning algorithms the prediction of important geochemical parameters from wireline logs can be satisfactorily achieved. Moreover, the results showed that samples with excellent hydrocarbon potential (high HI with low OI) are largely controlled by depositional environment. High hydrocarbon potential samples show an abundance of organic-rich lamalginites and filamentous alginite related to deeper offshore marine depositional environments. Low hydrocarbon potential samples (low HI and high OI) are related to the dilution by clastic terrigenous organic matter input due to proximity to deltaic sediment source. This is evident by the observed abundance of reworked vitrinite, and inertinite macerals scattered organic matter in the silty argillaceous matrix.

Fig. 10. Final Random Forest output of HI and OI prediction for Well A, Panther P-52, South Merasheen K-55 and South Tempest G-88 wells.

Fig. 10. (*continued*).

Fig. 11. Pseudo van Krevelen plot showing good correlation between measured HI and OI vs. Predicted HI vs. OI.

Contribution of authors

1) John B. Gordon (corresponding author)

- The conception and design of the study
- Analysis and interpretation of data (organic petrology, geochemistry)
- Drafting the article or revising it critically for important intellectual content
- 2) Hamed Sanei
- Interpretation of data (organic petrology, geochemistry)
- Revising the manuscript critically for important intellectual content

3) Per K. Pedersen

- Interpretation of data (depositional setting)
- Revising the manuscript critically for important intellectual content

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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